## COMMENTS

## Comment on "Are Vibrationally Excited Molecules a Clue for the O<sub>3</sub> Deficit Problem and HO<sub>x</sub> Dilemma in the Middle Atmosphere?"

Gregory P. Smith\* and Richard A. Copeland

Molecular Physics Laboratory, SRI International, Menlo Park, California 94025

Received: August 18, 2004

Recent measurements<sup>1</sup> of OH in the mesosphere are overpredicted by standard photochemical models. Coupled with underpredictions of data on HO<sub>2</sub> and O<sub>3</sub>, this predictive inconsistency has led to a reexamination of the relatively simple H–O kinetics in the upper atmosphere.<sup>2,3</sup> Varandas<sup>4</sup> has proposed that reactions involving vibrationally excited O<sub>2</sub> and OH can explain this situation, on the basis of what is reported to be a steady-state analysis of rates derived from detailed trajectory calculations.<sup>5</sup> We wish to point out several problems with this analysis.

The predicted concentrations of vibrationally excited OH, OH(v'), in Figure 6 of the Varandas manuscript<sup>4</sup> unexpectedly equal or exceed those for thermal  $OH^{6-8}$  ( $\sim 3 \times 10^6$  cm<sup>-3</sup>) and show an altitude dependence significantly different from rocket and satellite measurements of these vibrationally excited molecules. The altitude profile of vibrationally excited OH in the Earth's atmosphere is well-known through measurements of the emission from these levels in the Meinel bands.<sup>9</sup> The many mesospheric measurements and accompanying kinetic models of such densities,<sup>10-13</sup> none of which is referenced or discussed in the paper, invalidate both this altitude distribution and the amounts of OH(v') by orders of magnitude. Observed maximum OH( $v' \ge 3$ ) at 85 km is  $<10^4$  cm<sup>-3</sup>, not  $\sim 10^7$  as seen in Figure 6 near 80 km. Significant amounts are found at 80-95 km, not 60-80 km shown in Figure 6. The source of this OH(v') error lies in using H-atom concentrations taken from a textbook graph<sup>14</sup> that are over 2 orders of magnitude too large. Measured H amounts,<sup>15</sup> empirical and photochemical transport models,6-8,16,17 and review articles6,10 all point to an H-atom concentration at 75 km near  $2 \times 10^7$  cm<sup>-3</sup> rather than the  $10^{10}$ of Varandas<sup>4</sup> and do not reach amounts above  $2 \times 10^8$  at higher altitude. This error directly propagates into overstating any excited-state concentrations or kinetics effects.

Also, purported steady-state analyses are presented without detailed derivations. It appears the algebra is applied to a selected sequence or envisioned cycle of reaction steps, whereas proper analysis requires full consideration of all competing steps. More critically, multipliers are used without justification, for example applying the H + O<sub>3</sub> reaction rate 5 times in generating the OH(v') expression (eq 39 and following). The net result is that the HO<sub>x</sub> steady-state values relative to H are overstated by 5. This same general criticism may be applied to computations of the fractional effect of the enhanced mechanism on OH concentrations later in the paper.

 TABLE 1: Rate Parameter Additions for the Vibrationally

 Excited Mechanism of Varandas<sup>4</sup> <sup>a</sup>

reaction ( $k = A e^{-E/RT}$ )	$A (cm^3/molec/s)$	E (cal/mole)
$\overline{\mathbf{H} + \mathbf{O}_3 \rightarrow \mathbf{OH}(v') + \mathbf{O}_2}$	$1.40 \times 10^{-10}$	934.
$OH + O_3 \rightarrow H + O_2 + O_2(v)$ [20%]	$1.20 \times 10^{-12}$	1749.
$O_3 + h\nu \rightarrow O + O + O_2(\nu'')$ [73 km]	$9.75 \times 10^{-4/s}$	
$O_3 + h\nu \rightarrow O(^1D) + O_2(\nu)$ [73 km]	$7.35 \times 10^{-3}$ /s	
$OH(v') + O_3 \rightarrow HO_2(v) + O_2$	$1.70 \times 10^{-11}$	
$OH(v') + O_3 \rightarrow H + O_2 + O_2$	$2.24 \times 10^{-11}$	
$OH(v') + M \rightarrow OH + M$	$2.43 \times 10^{-11}$	
$HO_2(v) + O_3 \rightarrow OH + 2O_2$	$2.00 \times 10^{-14}$	1351.
$HO_2(v) + M \rightarrow HO_2 + M$	$2.43 \times 10^{-11}$	
$OH(v') + O_2(v'') \rightarrow H + O + O_2$	$8.00 \times 10^{-11}$	
$O_2(v'') + O_2(v) \rightarrow O_3 + O$	$8.00 \times 10^{-12}$	
$O_2(v'') + O_2 \rightarrow O_2(v) + O_2$	$3.00 \times 10^{-13}$	
$O_2(v) + O_2 \rightarrow 2O_2$	$4.00 \times 10^{-15}$	

<sup>*a*</sup> Base photolysis rates from Connell.<sup>18</sup> O<sub>2</sub> relaxation rates estimated from values summarized in Slanger and Copeland.<sup>21</sup>  $v'' \ge 25$ ;  $v' \ge 3$ ;  $v \ge 1$ .

Rather than descending into the complexities of further analysis, we have included the proposed mechanism modifications into box model simulations of the local photochemistry taken from outputs of the Lawrence Livermore National Laboratory 2-D diurnally varying model of the atmosphere.<sup>18</sup> This method is the simple and sure way to examine effects. The photolysis rates, kinetics rate constants,<sup>19</sup> species concentrations, and temperature are taken from the 2-D model output, and additional excited-state reactions are added. The selected locale is August 15 at noon, 50N latitude, for 43-, 55-, and 73km altitudes. The added steps and rate constants are shown in Table 1. Kinetics and photolysis rates were then integrated for 5 h using the Senkin code.<sup>20</sup> This time is in large excess of that required to achieve a photochemical steady state. The results of the box model simulation are given in Table 2. (Small changes from the initial concentrations reflect effects of temporally varying radiation and transport in the 2-D model.) No differences in final  $O_3$ , OH, or  $HO_2$  concentrations (<1%) were observed with the added excited-state kinetics. The proposed modifications have no effect, nullifying the conclusions reported from the Varandas analysis and casting doubt on the methodology.

In addition to the problems mentioned already, there are other substantive errors. The paper assumed that  $H + O_3$  is not just the main but the only source of OH (p 764), but in fact as most models show, the  $O(^{1}D) + H_{2}O$  reaction dominates below 50 km, and the rapid HO<sub>x</sub> recycling reaction  $O + HO_2 \rightarrow OH +$ O<sub>2</sub> must also be considered. Thus, an important OH-formation step is being ignored in this comparison of an excited-state production mechanism. Second, the "approach I" estimate of vibrational enhancement, which shows the larger effect, is based on eq 47. This equation erroneously double-counts the oxygen mole fraction in air (in both the 1/5 factor and the [O<sub>2</sub>] term). The value and enhancement are thus artificially inflated by a factor of 5. Third, no relaxation kinetics for vibrationally excited  $O_2$  can be found in the paper or its mechanisms. Such rate constants (mainly by O2) are known21 and necessary to determine the steady-state  $O_2(v)$  populations and rates; if excluded from the integrated kinetics modeling, wildly excessive

<sup>\*</sup> Corresponding author. E-mail: gregory.smith@sri.com.

 TABLE 2: Results of 5-Hour Photochemical Model Integration<sup>a</sup>

altitude	start O <sub>3</sub>	start OH	start HO <sub>2</sub>	reg. O <sub>3</sub>	reg. OH	reg. HO <sub>2</sub>	excit. O <sub>3</sub>	excit. OH	excit. HO <sub>2</sub>
km	ppm	ppb	ppb	ppm	ppb	ppb	ppm	ppb	ppb
73	0.167	4.05	1.85	0.186	3.74	1.70	0.186	3.74	1.70
55	1.46	0.867	0.403	1.55	0.827	0.383	1.55	0.827	0.383
43	4.56	0.298	0.190	4.66	0.288	0.185	4.66	0.289	0.184

<sup>a</sup> Last column results use Table 1 vibrationally excited mechanism additions.

amounts of  $O_3$  and  $HO_x$  result, especially at the lower altitudes. Finally, the rate constant given for k(33) on p 766 is 250 times too large.

We conclude from these facts and the direct simulations that the analysis which purports to show excited-state HO<sub>x</sub> chemistry as a potential solution to the ozone deficit and the HO<sub>x</sub> dilemma in upper atmosphere modeling is faulty, and demonstrate that the mechanism rate constants are insufficient to produce significant changes. Other searches for solutions will continue, but great care must be exercised in analyzing effects.

Acknowledgment. Research supported by the NASA ITM Physics Program (grants NAG5-11927, NAG5-10391, and NNG04GG05G) and the NSF Aeronomy Program (grants 99098078 and 0233523). Thanks to Dr. Peter Connell of LLNL for providing his model results.

## **References and Notes**

(1) Summers, M. E.; Conway, R. R.; Siskind, D. E.; Stevens, M. H.; Offermann, D.; Riese, M.; Preusse, P.; Stobel, D. F.; Russell, J. M., III *Science* **1997**, 277, 1967.

(2) Conway, R. R.; Summers, M. E.; Stevens, M. H.; Cardon, J. G.; Preusse, P.; Offermann, D. *Geophys. Res. Lett.* **2000**, *27*, 2613.

(3) Jucks, K. W.; Johnson, D. G.; Chance, K. V.; Traub, W. A.; Margitan, J. J.; Osterman, G. B.; Salawitch, R. J.; Sasano, Y. *Geophys. Res. Lett.* **1998**, *25*, 3935.

(4) Varandas, A. J. C. J. Phys. Chem. A 2004, 108, 758.

(5) Varandas, A. J. C. J. Phys. Chem. A 2003, 107, 3769.

(6) Roble, R. G. Energetics of the Mesosphere and Thermosphere. In The Upper Mesosphere and Lower Thermosphere: A Review of Experiment and Theory; Johnson, R. M., Killeen, T. L., Eds.; American Geophysical Union: Washington, DC, 1995; Vol. 87, p 1.

(7) Rodrigo, R.; Lopez-Moreno, J. J.; Lopez-Puertas, M.; Moreno, F.; Molina, A. *Planet. Space Sci.* **1986**, *34*, 723.

(8) Zhu, X.; Yee, J. H.; Lloyd, S. A.; Strobel, D. F. J. *Geophys. Res.* 1999, 104, 23995.

(9) Meinel, A. B. Astrophys. J. 1950, 111, 207.

(10) Meriwether, J. W., Jr. J. Geophys. Res. 1989, 94, 14629.

(11) Dodd, J. A.; Lipson, S. J.; Lowell, J. R.; Armstrong, P. S.;

Blumberg, W. A. M.; Nadile, R. M.; Alder-Golden, S. M.; Marinelli, W. J.; Holtzclaw, K. W.; Green, B. D. J. Geophys. Res. **1994**, 99, 3559.

(12) Adler-Golden, S. J. Geophys. Res. 1997, 102, 19969.

(13) Melo, S. M. L.; Lowe, R. P.; Russell, J. P. J. Geophys. Res. 2000, 105, 12397.

(14) Steinfeld, J. I.; Francisco, J. S.; Hase, W. L. Chemical Kinetics and Dynamics, 2nd ed.; Prentice Hall: Upper Saddle River, NJ, 1999.

(15) Sharp, W. E.; Kita, D. J. Geophys. Res. 1987, 92, 4319. Thomas,
 R. J. J. Geophys. Res. 1990, 95, 16457. Adler-Golden, S.; Gruninger, J.;
 Smith, D. R. J. Geophys. Res. 2992, 97, 19509.

(16) Hedin, A. E. J. Geophys. Res. 1991, 96, 1159.

(17) Picone, J. M.; Hedin, A. E.; Drob, D. P.; Aikin, A. C. J. Geophys. Res. 2002, 107 (A12), 1468.

(18) Connell, P. S. Private communication, 2002.

(19) Sander, S. P.; Friedl, R. R.; Golden, D. M.; Kurylo, M. J.; Huie, R. E.; Orkin, V. L.; Moortgat, G. K.; Ravishankara, A. R.; Kolb, C. E.; Molina, M. J.; Finlayson-Pitts, B. J. *Chemical Kinetics and Photochemical Data for Use in Stratospheric Modeling, Evaluation Number 14*; National Aeronautics and Space Administration, Jet Propulsion Laboratory: Pasadena, CA, 2003.

(20) Lutz, A. E.; Kee, R. J.; Miller, J. A. SENKIN: A Fortran Program for Predicting Homogeneous Gas-Phase Chemical Kinetics with Sensitivity Analysis; Sandia National Laboratories: Livermore, CA, 1988.

(21) Slanger, T. G.; Copeland, R. A. Chem. Rev. 2003, 103, 4731.